Global temperature is increasing, especially over northern lands (>50° N), owing to positive feedbacks. As this increase is most pronounced in winter, temperature seasonality (S_T)—conventionally defined as the difference between summer and winter temperatures—is diminishing over time, analogous to its equatorward decline at an annual scale. The initiation, termination, and performance of vegetation photosynthetic activity are tied to threshold temperatures. Trends in the timing of these thresholds and cumulative temperatures above them may alter vegetation productivity, or modify vegetation seasonality (S_R) over time. Therefore, the relationship between S_T and S_R is critically examined here with newly improved ground and satellite data sets. The observed diminishment of S_T and S_R is equivalent to 4° and 7° (5° and 6°) latitudinal shift equatorward during the past 30 years in the Arctic (boreal) region. Analysis of simulations from 17 state-of-the-art climate models indicates an additional S_T diminishment equivalent to a 20° equatorward shift this century. How S_R will change in response to such large projected S_T declines and the impact this will have on ecosystem services are not well understood, hence the need for continued monitoring of northern lands as their seasonal temperature profiles evolve to resemble those further south.

The Arctic (8.16 million km²) is defined here as the vegetated area north of 65° N, excluding crops and forests, but including the tundra south of 65° N. The boreal region (17.86 million km²) is defined as the vegetated area between 45° N and 65° N, excluding crops, tundra, broadleaf forests and grasslands south of the mixed forests, but including needleleaf forests north of 65° N (Supplementary Fig. S1). These definitions are a compromise between ecological and climatological conventions. Importantly, they include all non-cultivated vegetation types within these two regions.

Comparisons of changes in seasonality of physical and biological variables require definitions that are concordant, have an ecological underpinning, for example, vegetation photosynthetic activity in the north depends on the seasonal cycle of temperature and not on the difference between annual maximum and minimum temperatures, and satisfy the principle that seasonality increases with latitude at an annual timescale owing to patterns of insolation resulting from Sun–Earth geometry alone (Fig. 1a and Supplementary Information S2.A). Therefore, S_T is defined as \[ 1 \div \overline{T}_l(l) \], where \( \overline{T}_l(l) \) is the zonally averaged annual mean temperature at latitude \( l \). S_R is analogously defined as \[ 1 \div \overline{N}_l(l) \], where \( \overline{N}_l(l) \) is the zonal mean of photosynthetic activity averaged over the photosynthetically active period (PAP) at latitude \( l \). These definitions possess the above-mentioned attributes and accurately represent the respective seasonal cycles (Supplementary Information S2.A.3).

The latitudinal profiles of PAP-mean temperature from 50° N to 75° N (ice sheets excluded throughout) show warming of 1–2°C between the early 1980s and late 2000s (Fig. 1b). Analogous profiles of normalized difference vegetation index (NDVI), a proxy for vegetation photosynthetic activity, show a similar increase. S_T is tightly coupled to S_T in the north (Fig. 1c). The slope of this relationship (\( \beta_{VT} \)) has not changed in the past 30 years (Fig. 1c, inset). Figure 1b,c may thus indicate widespread and matching patterns of temperature and NDVI increase and corresponding reductions in S_T and S_R throughout northern lands. If this were to continue, significant increases in productivity may be expected in the boreal Arctic region during this century on the basis of climate model projections of large S_T diminishment (Fig. 4c), even as insolation seasonality remains unchanged, which would have major ecological, climatic and societal impacts. Therefore, the apparent constancy of \( \beta_{VT} \) in Fig. 1c is tested in four ways.

In the first test, the constancy of \( \beta_{VT} \) is based on widespread statistically significant increases in PAP-mean NDVI and temperature. This is assessed using four statistical models. Results from two statistically robust models are mainly discussed here (Models 3 and 4 in Supplementary Information S2.C.1).
Second, the greening is most prominently seen in coastal tundra\(^6\) and eastern mixed forests in North America, needleleaf and mixed forests in Eurasia, and shrublands and tundra in Russia (Fig. 2a and Supplementary Fig. S7). North American boreal vegetation shows a fragmented pattern of greening and browning\(^7\), unlike its counterpart in Eurasia, which shows widespread contiguous greening. Further analysis reveals little evidence of widespread browning of boreal vegetation at the circumpolar scale (Supplementary Information S3.A).

Third, about 90% of the Arctic and 70% of the boreal greening vegetation show \(N_p\) increases >2.5% per decade (Fig. 2c). These changes in \(N_p\) can be expressed as changes in PAP duration. For example, a trend of +x days per decade at a location in Fig. 2b means that the vegetation there would require x more days of PAP in 1982, the first year of the NDVI record, to equal its \(N_p\) ten years later. About 88% of the Arctic and 81% of the boreal greening vegetation show extensions in PAP >3 days per decade (Fig. 2d). These extensions hint of \(S_V\) declines in these two regions—this is further explored in the fourth test below.

Next, regarding temperature changes, PAP-mean temperature could not be accurately evaluated because of the coarse temporal resolution of temperature data (monthly). Therefore, statistical analysis was performed on a per-pixel basis but using a close analogue, May–September (warm-season) average temperature, \(T_{WS}\). The proportion of Arctic and boreal regions exhibiting statistically significant increase in \(T_{WS}\) varied from 51 to 54% (Supplementary Table S4 under the heading Significant Trends; Supplementary Fig. S8). The proportion exhibiting statistically significant decrease in \(T_{WS}\) was <0.6%.

Therefore, the constancy of \(\beta_{VT}\) is based on widespread statistically significant increases in PAP-mean NDVI (34–41%) and its temperature analogue \(T_{WS}\) (51–54%) in the study area.

In the second test, the constancy of \(\beta_{VT}\) is based on spatially matching statistically significant changes in \(N_p\) and \(T_{WS}\). The sign of significant trends in \(N_p\) and \(T_{WS}\), or lack of such trends, is similar in about 47% of the Arctic and boreal vegetated lands (Fig. 3a,b; all model results in Supplementary Fig. S9 and Supplementary Table S4). The trends of \(N_p\) and \(T_{WS}\) are of opposite sign in <2% of the study area. Greening or browning is not observed in an additional 27–31% of vegetated lands where warming is moderate. This pattern is seen in evergreen needleleaf forests of eastern North America, deciduous needleleaf forests of Russia and in patches in western Canada and Alaska. Thus, in nearly 74–78% of the Arctic and boreal regions, trends in \(N_p\) and \(T_{WS}\) did not strongly oppose one another during the past 30 years. Therefore, the constancy of \(\beta_{VT}\) is based on spatially matching statistically significant changes in \(N_p\) and \(T_{WS}\).

In the third test, \(\beta_{VT}\) is spatially invariant, that is, coefficients \(\beta_{VT}\) of the Arctic and boreal region are similar. Statistical analysis with two regression models\(^8\) indicates highly significant (\(p < 0.01\)) relationships between \(S_T\) and \(S_V\) anomaly time series in both regions (Fig. 3c,d and Supplementary Table S5). Here, \(S_V\) is defined in terms of PAP-mean temperature for large zonal bands such that it satisfies the Sun–Earth geometric definition of seasonality. The coefficients associated with the temperature variable of the two regions are statistically similar in both models. Therefore, \(\beta_{VT}\) is spatially invariant over the 30-year study period.

In the fourth test, \(\beta_{VT}\) is spatially and temporally invariant, that is, coefficients \(\beta_{VT}\) of the Arctic and boreal regions are not only similar but also did not change between the first and second halves of the 30-year study period. To avoid performing statistical analysis on short data records, changes in \(S_T\) and \(S_V\) were translated into latitudinal shifts during each half of the study period and compared with one another. Briefly, data from the early part of the time series were used to define baselines depicting seasonality variation with respect to latitude in the Arctic and boreal regions. The location of
temperature and vegetation seasonality on the respective baselines for three periods yielded seasonality declines in terms of latitude between the first half (mid 1990s and early 1980s) and second half (late 2000s and mid 1990s) of the data record.

The early-1980s (1982–1986) Arctic warm-season $S_T$, corresponding to the warm-season $S_T$ of vegetated lands $>$64.8° N (Fig. 4a). By the late 2000s, the warm-season temperature profile of the Arctic was similar to the early-1980s warm-season temperature profile of vegetated lands $>$60.8° N—a decline in $S_T$ of 4.0° in latitude. The early-1980s boreal region warm-season $S_T$ corresponded to the warm-season $S_T$ of vegetated lands between 45° N and 66.1° N. By the late 2000s, the warm-season temperature profile of the boreal region was similar to the early-1980s warm-season temperature profile of vegetated lands between 45° N and 60.9° N—a decline in $S_T$ of 5.2° in latitude. Changes in $S_T$ were similarly quantified (Fig. 4b). The corresponding declines in Arctic and boreal $S_T$ are 7.1° and 6.3° in latitude.

The difference in $S_T$ decline between the first and second halves of the 30-year period is negligible in both the Arctic and boreal region, in view of the coarse resolution of temperature data. However, this is not the case with $S_V$. The Arctic $S_V$ decline accelerated, that is, the greening rate increased over time, from 2.15° latitude between the early 1980s and mid 1990s to 4.9° latitude between the mid 1990s and late 2000s. In contrast, $S_V$ decline in the boreal region decelerated from 5.7° to 0.6° latitude. These varying rates of $S_V$ declines are inconsistent with the idea of a spatially and temporally invariant $\beta_{VT}$.

In summary, the first three tests support the observed (Fig. 1c) tight coupling between $S_V$ and $S_T$. However, the fourth test indicates that $\beta_{VT}$ varies with time and that this variation differs between the Arctic and boreal regions, with greening in the Arctic accelerating over time, whereas boreal greening is decelerating over time. The robustness of these conclusions is addressed in Supplementary Information S3.B.
Empirical evidence suggests that in addition to direct effects of warming\textsuperscript{11,12} several other factors influence $\beta_{VT}$ (refs 13--15). These include: warming-induced disturbances and recovery (summertime droughts\textsuperscript{16}, mid-winter thaws\textsuperscript{17}, increased frequency of fires and outbreaks of pests\textsuperscript{18}, shrinking and draining of lakes from thawing permafrost\textsuperscript{19}, desiccation of ponds\textsuperscript{20}, colonization of the growing banks by vegetation\textsuperscript{21} and so on), interacting effects of temperature and precipitation\textsuperscript{22}, complex feedbacks (feedbacks that enhance wintertime snow amount on land asymmetrically between Eurasia and North America\textsuperscript{23}, feedbacks from declining snow-cover extent on land\textsuperscript{1} leading to longer growing seasons\textsuperscript{3,9} and promoting vegetation compositional/structural changes\textsuperscript{12,13,24,25}, enhanced nitrogen mineralization in warmer soils insulated by increased shrub cover\textsuperscript{26} and so on), anthropogenic influences (pollution from metal smelters\textsuperscript{27}, herding practices of grazing herbivores\textsuperscript{28} and so on) and changes in wild herbivore populations\textsuperscript{29}. These factors could have contributed to an amplification of $\beta_{VT}$ in the Arctic and dampening in the boreal region.

Projections of $S_T$ changes during this century are of interest given the observed relationship between $S_T$ and $S_V$ of the past 30 years. The median decline $S_T$ in the Arctic and boreal regions from 17 climate models is 22.5° and 21.8° latitude by the decade 2091--2099.
By 2091–2099, the annual temperature profile of the Arctic (boreal) is projected to be similar to the baseperiod annual temperature profile of lands north of 42.4° N (23.4° N).

The observed decline during 2001–2010 is already greater than the multi-model median estimate (Supplementary Table S6). Recent trends are thus consistent with longer-term observations. It is not known how \( S_\ell \) will change in response to large projected declines in \( S_\ell \) as this depends on adaptability of extant species and migration rates of productive southerly species in the face of unchanging insolation seasonality, increased frequency of winter warming events and other factors (Supplementary Information S3.C), hence the need for continued monitoring of northern lands as their seasonal temperature profiles evolve to resemble those further south.

### Methods

All satellite and ground data used in this research are described in Supplementary Information S1. The derivation, testing and justification of temperature and vegetation seasonality definitions are described in Supplementary Information S2.A. The method for estimation of PAP is described in Supplementary Information S2.B. The four statistical methods employed to assess statistical significance and magnitude of trends are described in Supplementary Information S2.C. The evaluation of temperature and vegetation seasonality baselines and diminishment over time are described in Supplementary Information S2.D–S2.G.

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### References


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The analysis was performed by X.L., R.B.M, Z.Z and J.B. All authors contributed with ideas, writing and discussions.

Additional information
Supplementary information is available in the online version of the paper. Reprints and permissions information is available online at www.nature.com/reprints. Correspondence and requests for materials should be addressed to X.L or R.B.M.

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The authors declare no competing financial interests.
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